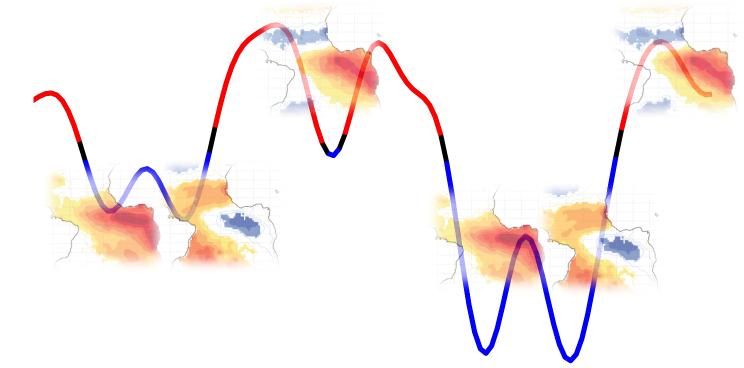
On the inter-annual tropical Atlantic variability modes under AMO phases in the observational record

(Under review in Journal of Climate)





University of **Reading**

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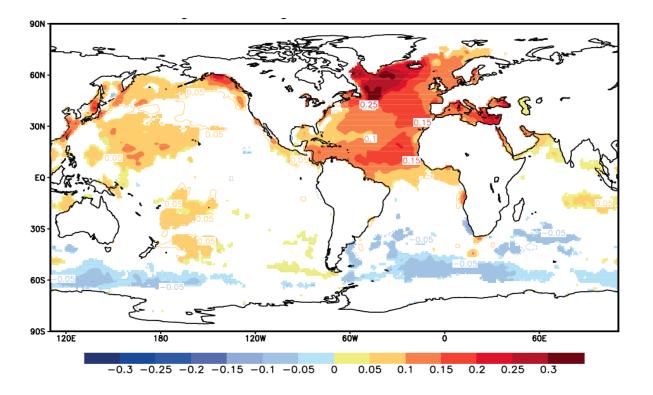
Teresa Losada⁽¹⁾ and Alban Lazar ⁽²⁾

(1) UCM, Madrid, Spain; (2) LOCEAN-IPSL, UPMC, France; (3) University of Reading, UK



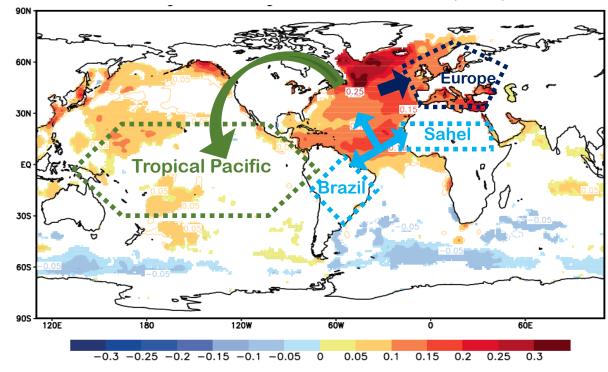






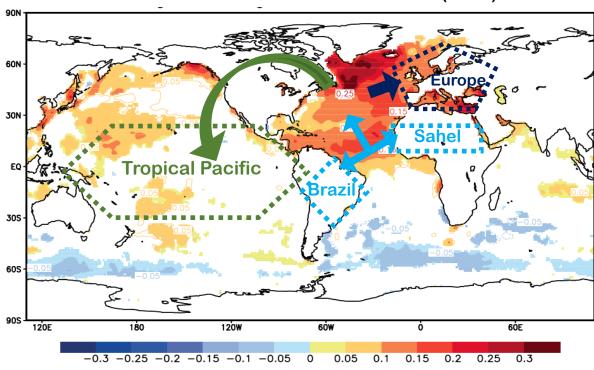
The decadal internal SST variability in the Atlantic Ocean is driven by the Atlantic Multidecadal Oscillation, AMO (Kerr 2000; Knight et al. 2006)

Atlantic Multidecadal Oscillation (AMO)



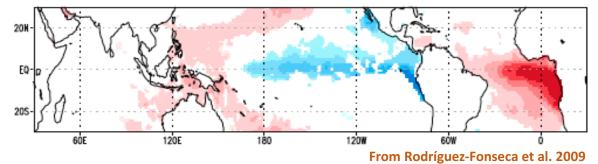
The AMO seems to modulate:

- The Tropical Pacific mean state and ENSO variability (Dong et al. 2006; Dong and Sutton 2007; Timmerman et al. 2007).
- The Brazilian and Sahelian precipitation (Knight et al. 2006).
- European summer climate (Sutton and Hodson 2003; Sutton and Dong 2012).



Atlantic Multidecadal Oscillation (AMO)

Atlantic-Pacific Niños connection



The Atlantic Niño teleconnections have been changed and strengthened after the 1970s (Polo 2008; Kucharski et al. 2009; Rodríguez-Fonseca et al. 2009; Ding et al. 2012; Losada et al. 2012a; Losada and Rodríguez-Fonseca 2015).

The multidecadal modulation of the Atlantic-Pacific Niños connection could be driven by AMO (Martín-Rey et al. 2014).

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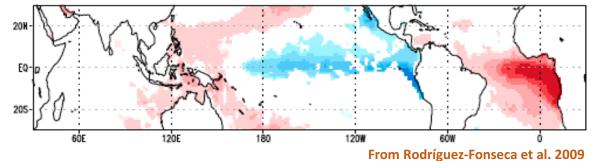
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Atlantic Multidecadal Oscillation (AMO) 60N · 30N **Tropical Pacific** 30S -60S · 90S 120E 180 1200 6ÓW 6ÖE -0.3 -0.25 -0.2 -0.15 -0.1 -0.05 0.15 0.2 0.25 0 0.05 0.1

The AMO seems to modulate:

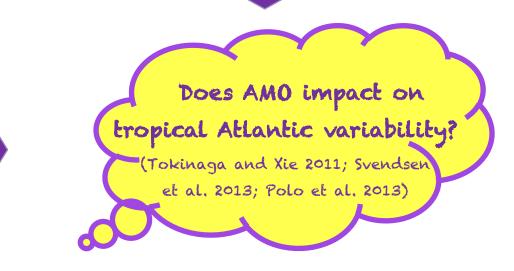
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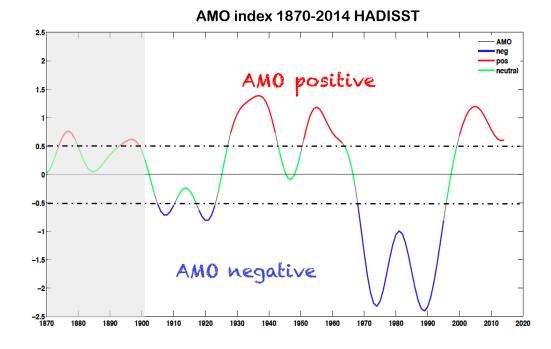
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DATA AND METHODS

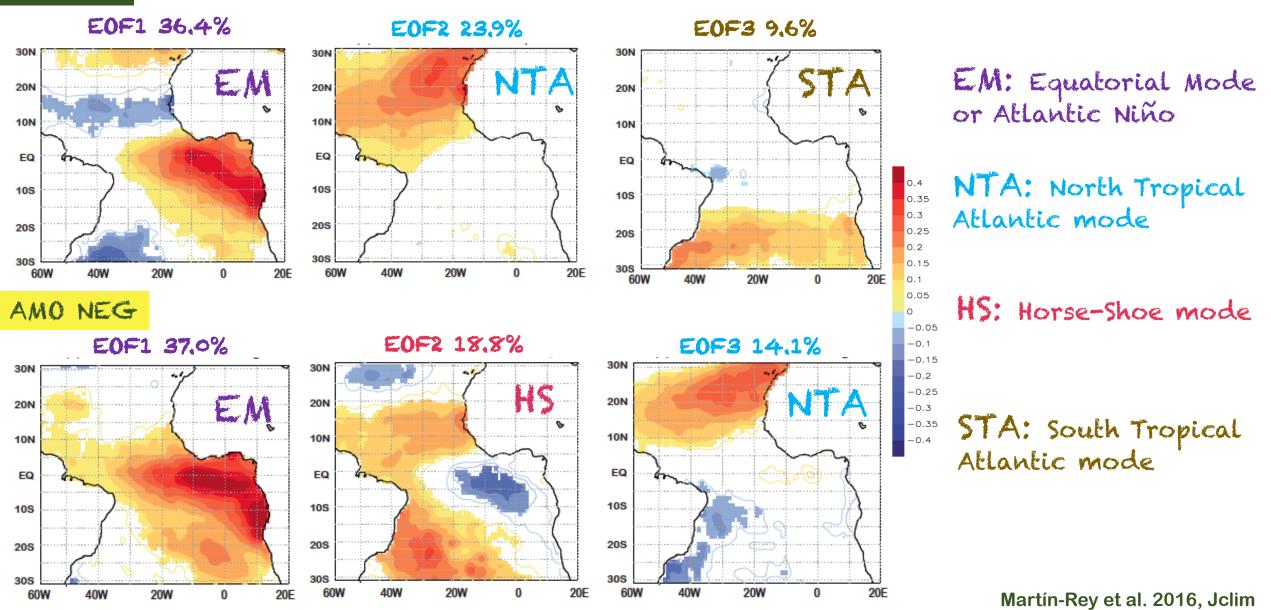
- * SST for 1870-2014 from HADISST (Rayner et al. 2003)
- * Surface WIND and SLP for 1900-2010 from ERA20C reanalysis (Poli et al. 2013)
- * Thermocline depth (Isotherm of 16C, D16) for 1871-2008 from SODA reanalysis (Giese and Ray 2011)
- * High-Pass (7 year cutt-of) Butterworth filter (Butterworth 1930)
- * Common period of study: 1900-2008



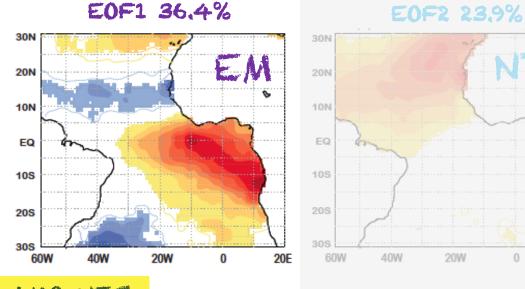
- * Principal Component Analysis (PCA) applied to the summer (JJAS) SST anomalies for different AMO phases:
- ★ 34 years for positive AMO phase: 1928, 1929, 1939, 1931, 1932, 1933, 1934, 1935, 1936, 1937, 1938, 1939, 1940, 1941, 1942, 1951, 1952, 1953, 1954, 1955, 1956, 1957, 1958, 1959, 1960, 1999, 2000, 2001, 2002, 2003, 2004, 2005, 2006, 2007.
- So years for negative AMO phase 1906, 1907, 1908, 1909, 1918, 1919, 1920, 1921, 1922, 1923, 1968, 1969, 1970, 1971, 1972, 1973, 1973, 1974, 1975, 1976, 1977, 1978, 1979, 1980, 1981, 1982, 1983, 1984, 1985, 1986, 1987, 1988, 1989, 1990, 1991, 1992, 1993, 1994, 1995

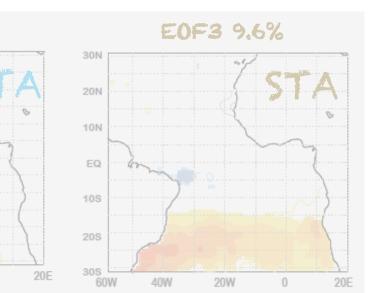
RESULTS: Does the TA SST variability change under AMO phases?

AMO POS



AMO POS

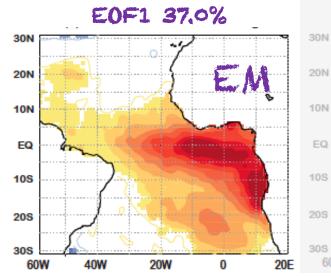




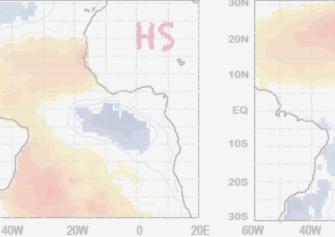
NEGATIVE AMO:

Atlantic Niño show larger amplitude and fraction of explained variance. The warm tongue presents a westward extension

AMO NEG







E0F3 14.1%

20W

0

EQ

10S

20S

305

60W

40W

20W

0

20E

20E

0

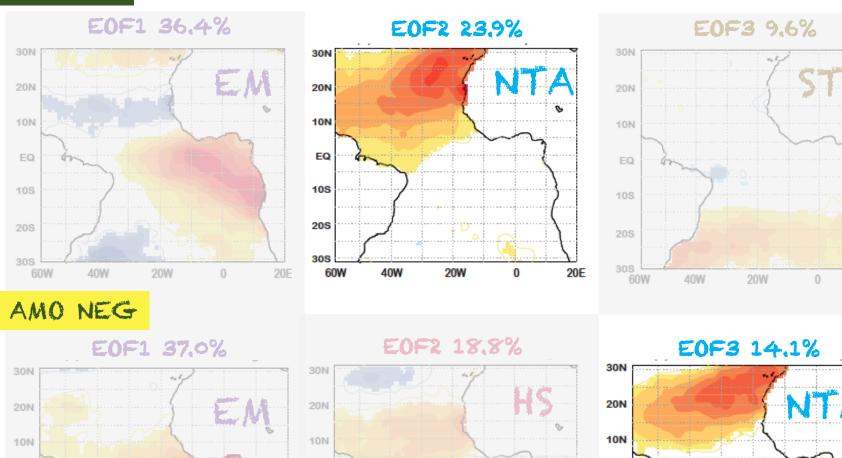
AMO POS

EQ

10S

205

60W



EQ

10S

20S

20E

40W

20W

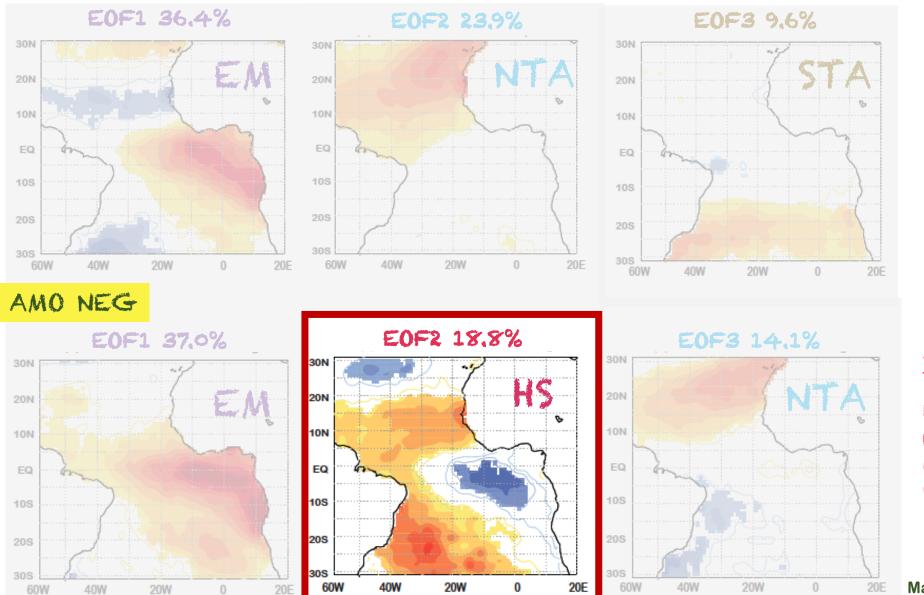
NEGATIVE AMO:

20E

NTA mode is relegated to the third mode and shows a meridional northward displacement

Martín-Rey et al. 2016, Jclim

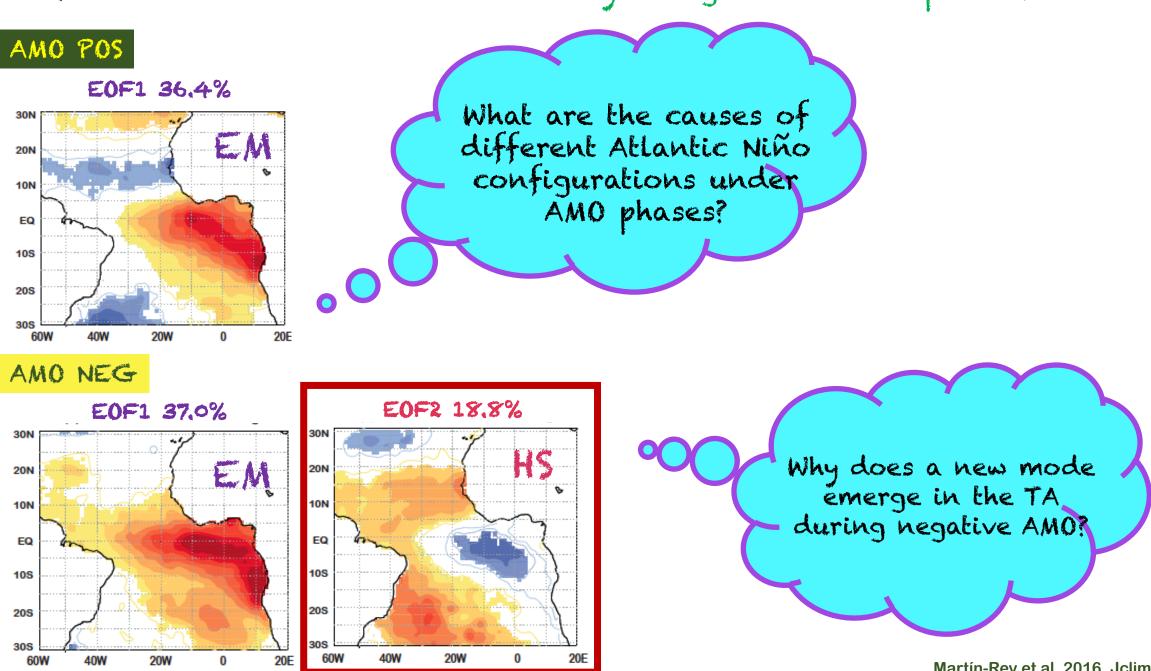
AMO POS



NEGATIVE AMO:

HS mode only emerges in negative AMO periods as the second variability mode.

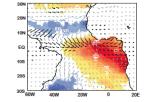
Martín-Rey et al. 2016, Jclim (under review)



Martín-Rey et al. 2016, Jclim (under review)

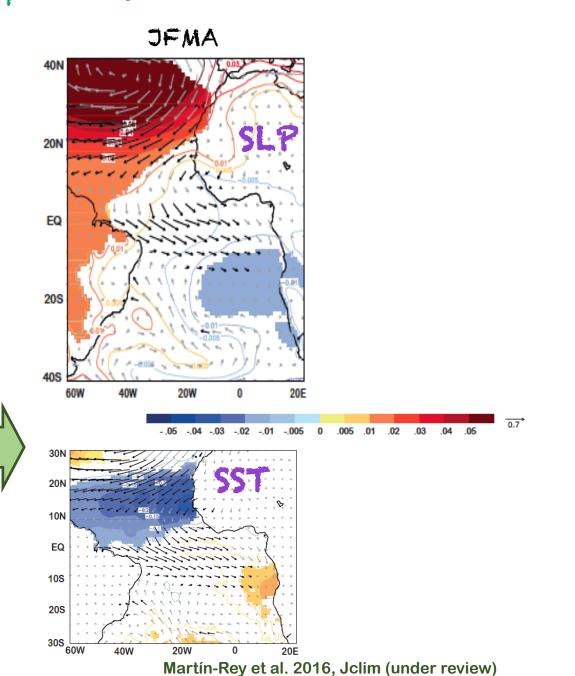


ATLANTIC NIÑO



JFMA:

- > Strengthening of Azores High → intensified north-easterlies → enhanced the evaporative cooling.
- > Low pressure in eastern STA → weakened southern trades → reduced the AB upwelling
 *Local winds (Polo et al. 2008; Richter et al. 2010)
 *Wind-excited equatorial KW (Lübbecke et al. 2010)



0.4 0.35 0.3 0.25

0.2 0.15

0.1 0.05

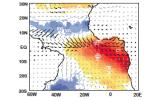
-0.15 -0.2

-0.25 -0.3 -0.35

-0.4



ATLANTIC NIÑO

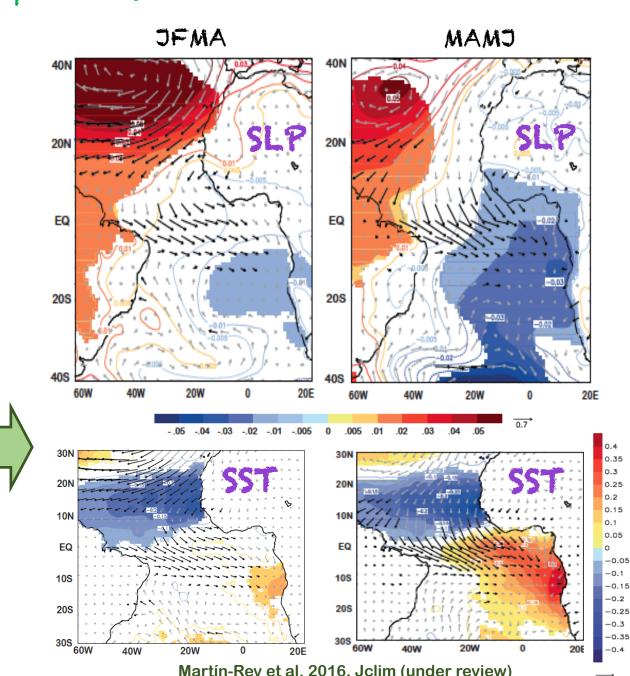


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MAMJ:

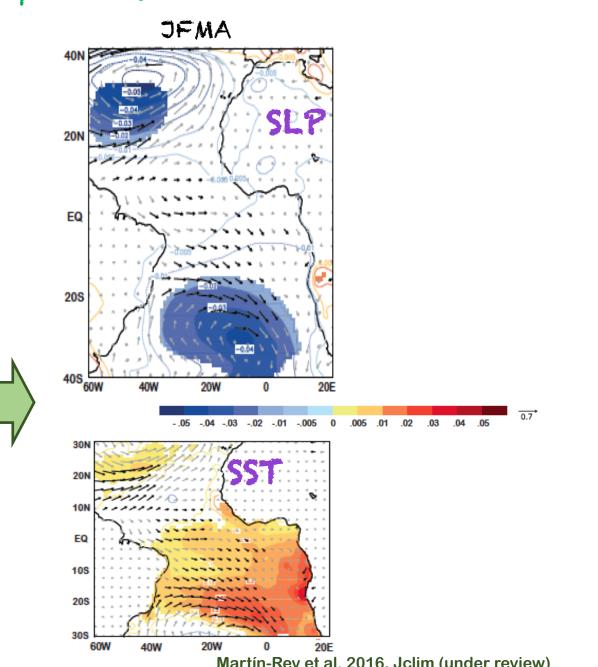
- Meridional evolved into east-west tropical SLP
 gradient → reinforced westerlies at the equator
 → deepened thermocline and warming.
 * Bjerknes feedback (Keenlyside and latif 2007)
 *Wind-excited equatorial KW (Lübbecke et al. 2010)
- > Negative Meridional Mode-type SST pattern in spring precedes the Atlantic Niño (servain et al. 1999; Foltz et al. 2010;Lübbecke and McPhaden 2012).





JFMA:

- Simultaneous weakening of Azores and St Helena High → reduction north-eastern and south-eastern trades → Decreased Latent heat Loss →warm the underneath regions.
- > Anomalous southward alongshore winds in AB region or excited KW → deepened thermocline → reduced the upwelling (Polo et al. 2008; Richter et al. 2010 al. 2010; Lübbecke et al. 2010).



0.4 0.35 0.3

0.25

0.2 0.15

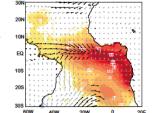
0.1 0.05

> -0.05 -0.1

> -0.15 -0.2

-0.25 -0.3 -0.35



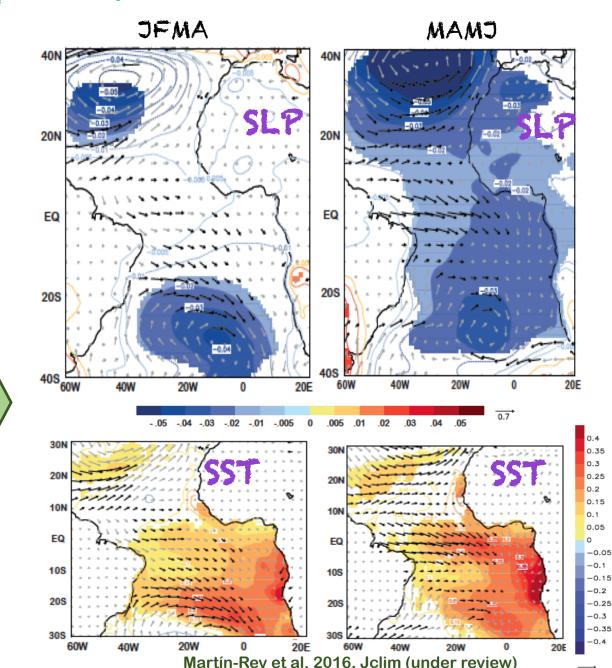


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MAMJ:

- > Large scale anomalous low SLP conditions \rightarrow general reduction of the tropical trades.
- > Stronger westerlies along the equatorial band → deepened thermocline → equatorial and coastal warming → Bjerknes feedback (Keenlyside and latif 2007)

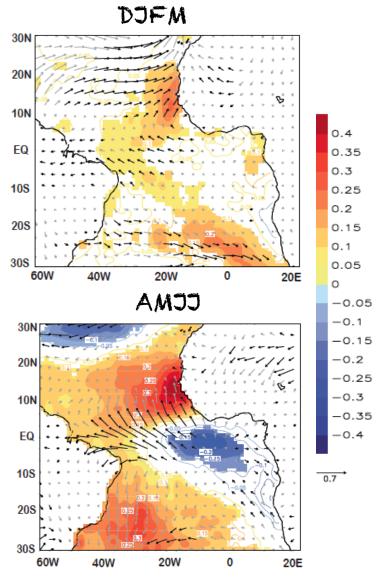


RESULTS: Which are the atmospheric forcings of the Horse-Shoe

AMO NEG

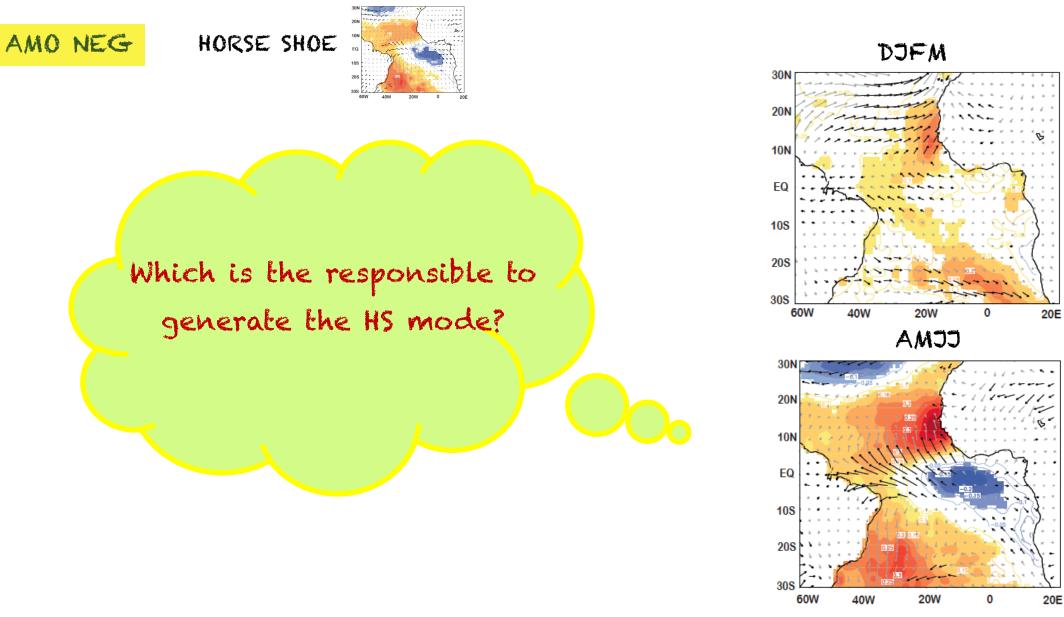
HORSE SHOE

- > Weakened northern trades and south-easterly trades → reduced evaporative cooling.
- > Anomalous northward winds along Senegal-Mauritanian coast → deepen thermocline through Ekman pumping and reduce the upwelling.
- \succ Meridional SST gradient between NTA and EQ \rightarrow reinforced cross-equatorial winds
- > Intensified equatorial easterly winds → enhanced eastern equatorial upwelling.
- ➤ The cold tongue developed → shallows the thermocline and reinforced the easterlies → activating the Bjerknes feedback (Keenlyside and latif 2007).



Martín-Rey et al. 2016, Jclim (under review)

RESULTS: Which are the atmospheric forcings of the Horse-Shoe



Martín-Rey et al. 2016, Jclim (under review)

0.4

0.35 0.3

0.25 0.2

0.15

0.05

-0.05 -0.1 -0.15 -0.2

-0.25 -0.3

-0.35

-0.4

0.7

0

RESULTS: Is the HS mode remotely forced?

Regression HS - SSTglob JFMA Regression HS - SSTglob MAMJ 90N 90N 60N 60N El Niño pattern during previous winter forced the HS mode. 30N EQ EQ 305 305 60S 60S 90S 90S 120E 180 120W 60W 60E 120E 180 120W 60W 60E 0 .15 .2 .25 .3 -.3 -.25 -.2 -.15 -.1 .05 -.05 0 .1 **Regression HS – SLPWIND glob JFMA Regression HS – SLPWIND glob MAMJ** Atmospheric Rossby 601 waves: PSA and PNÅ patterns (Handoh et al. 2006a,b; 30N López-Parages and Rodríguez-Fonseca 2012) 30S Walker circulation (wang 2006) 60S 90S 120E 90S 60E 180 120W 60W 120F 60W 180 120W

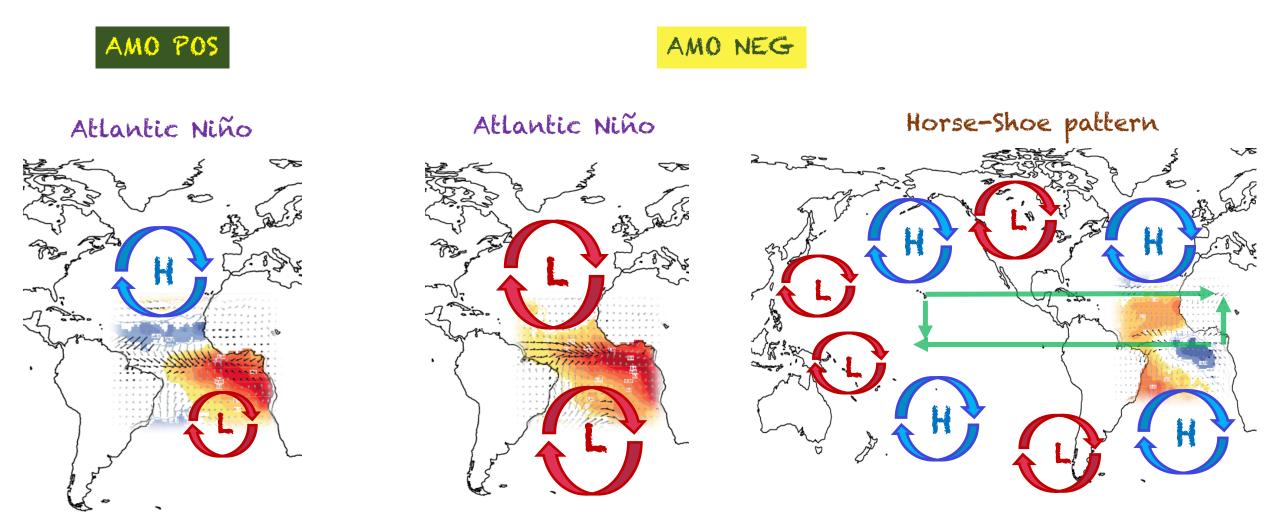
Martín-Rey et al. 2016, Jclim (under review)

-.07 -.06 -.05 -.04 -.03 -.02 0 .02 .03 .04 .05 .06 .07

0.7

Contrasting with previous results (Polo et al. 2008; Lübbecke et al. 2010; Mamchi et al. 2016):

Both Subtropical High Pressure Systems seems to drive the development of different Atlantic Niño configurations and HS pattern under AMO phases.



CONCLUSIONS

> Observed inter-annual TAV seems to be modulated by AMO: the emergence and spatial configuration of the modes vary under AMO phases.

> During negative AMO phases the Atlantic Niño presents larger amplitude and a westward extension of its warm tongue. Its development is associated with the weakening of both Azores and St Helena Highs, reducing the tropical trades.

> For positive AMO phases, the Atlantic Niño is related to a meridional SLP pattern, intensifying the northern trades. This SLP configuration evolves into a east-west tropical gradient, reinforcing the westerlies along the equatorial Atlantic.

CONCLUSIONS

> A new TAV pattern shows up during negative AMO phases: the Horse-Shoe mode. It is characterized by a horse-shoe of positive SST anomalies in NTA and STA, surrounding an anomalous cooling in the eastern equatorial Atlantic.

> HS only emerges during negative AMO phases, forced by an El Niño phenomenon from previous winter.

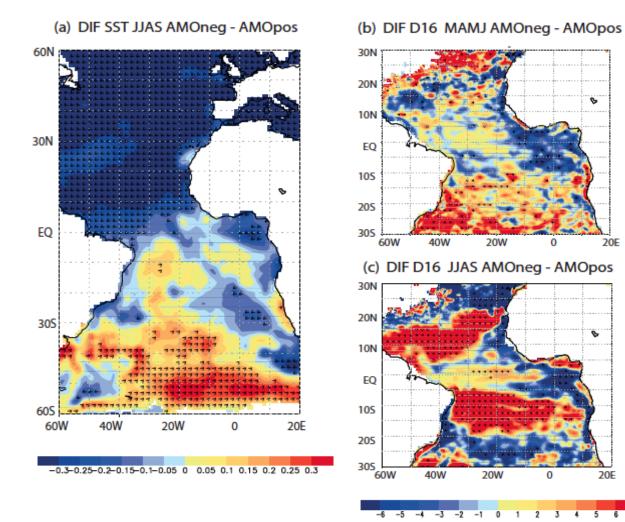
> These changes in Atlantic Niño configuration and HS pattern are related to the interaction of both Subtropical Highs Pressure Systems.

THANK YOU FOR YOUR ATTENTION!



The research leading to these results received funding from the EU FP7/2007-2013 under grant agreement no.603521 (PREFACE project). This study was also supported by the Spanish MINNECO project CGL2012-38923-C02-01.

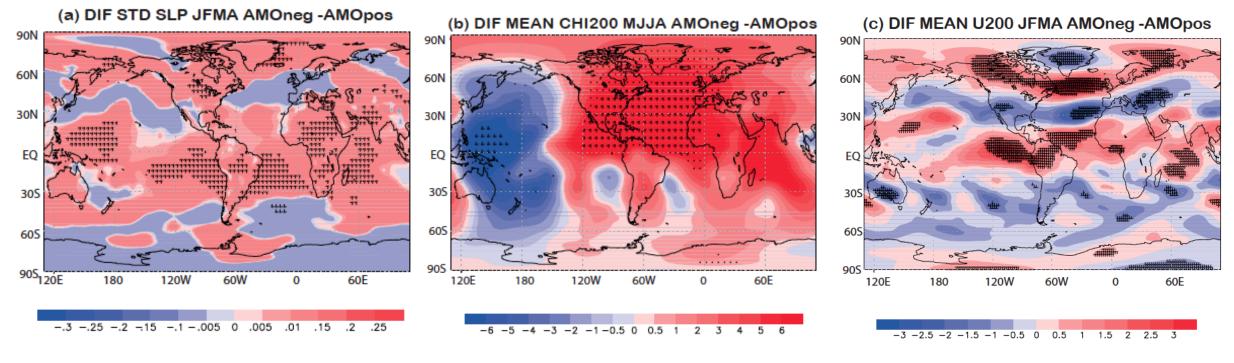
> Changes in the Atlantic Niño amplitude could be caused by shallower thermocline along the equatorial band from boreal spring to summer months, which increases the effect of the Bjerknes feedback and in turn the SST variability during negative AMO phases.



20E

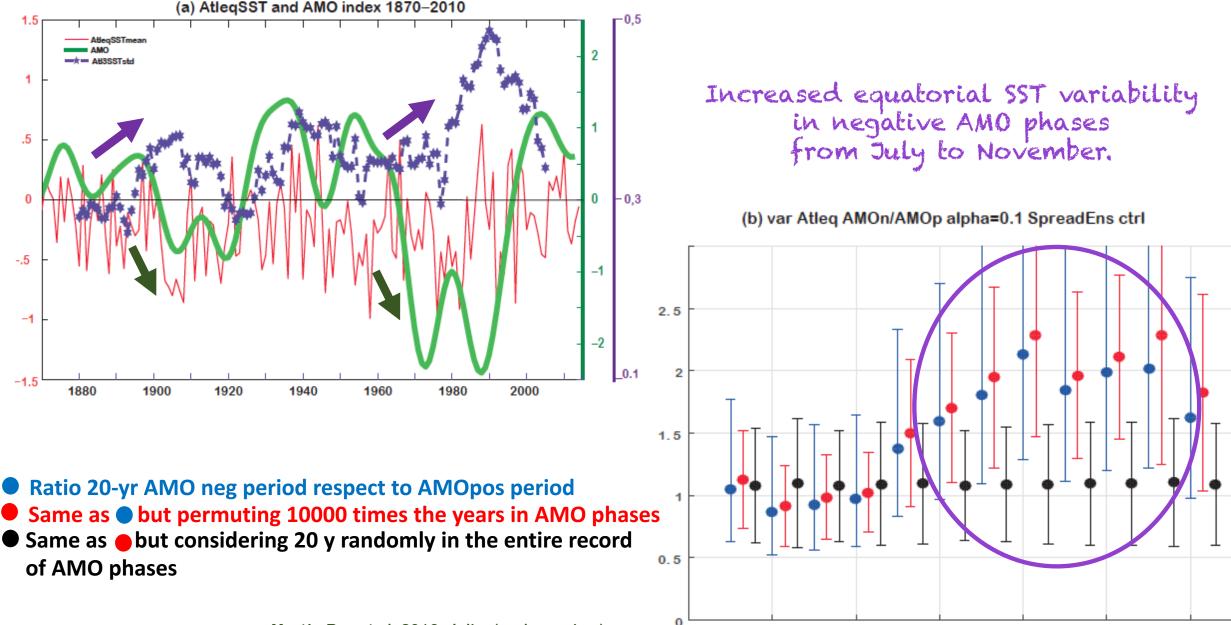
20E

> An enhancement of ENSO tropical-extratropical teleconnections over the Atlantic, together with an intensification of the Walker circulation could favour the ENSO impact on the TAV during negative AMO periods.



Martín-Rey et al. 2016, Jclim (under review)

CHANGES IN THE EQUATORIAL ATLANTIC SST VARIABILITY ALONG THE 20th CENTURY



JAN

FEB

MAR APR MAY

JUN JUL

AUG SEP

OCT NOV

DEC

